

**ERTH 456 / GEOL 556**  
**Volcanology**

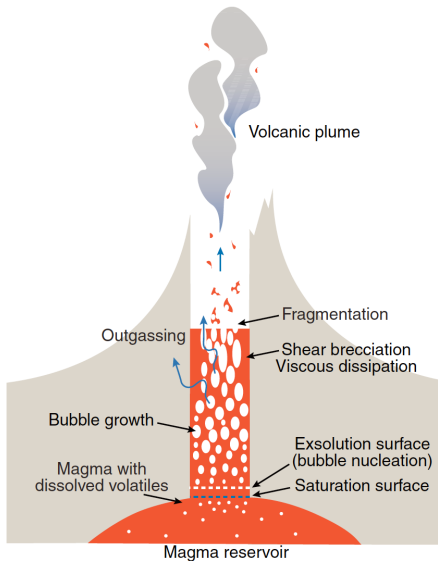
**– Lecture 08: Explosive Eruptions –**

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hours: TR 3-4PM or appt.

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What happens once magma makes it to the surface?

# Volcano Anatomy



*Gonnermann & Manga, 2007*

## Quiz: Why doesn't this work?

`https://youtu.be/g7Ovo9oh11I?t=41s`

# How do we get an explosive eruption?

- fragmentation at top of magma column (don't lose all gases through conduit wall)
- sufficient flow velocity
- vent geometry



USGS, 2016; Framepool, 2011

# Makeup of Explosive Eruptions

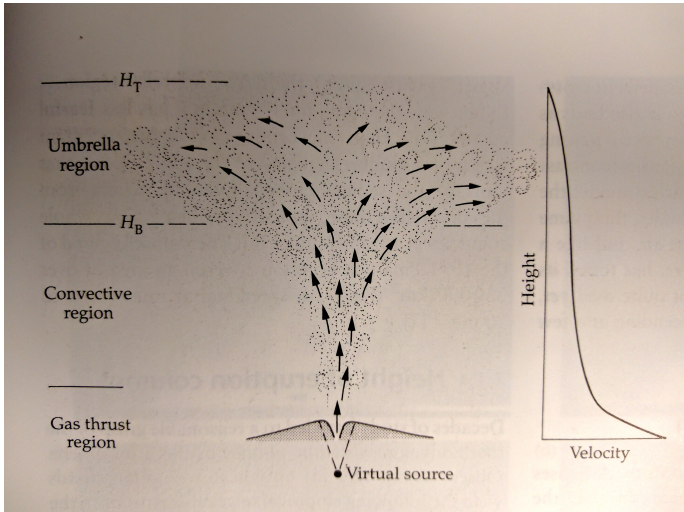
- volcanic plumes
- pyroclastic fall deposits
- pyroclastic density currents

# Makeup of Explosive Eruptions



*ISS Crew, 2008*

# Plume Basics



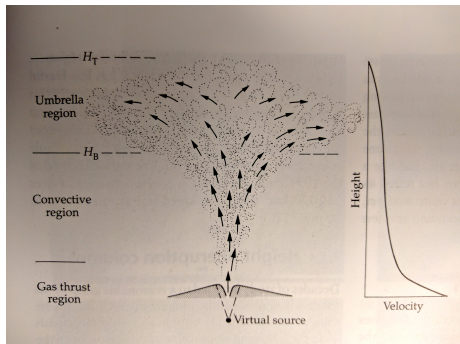
*Francis & Oppenheimer, 2004*

$H_B$ : height of neutral buoyancy;  $H_T$ : maximum plume height  
( $H_T \approx 1.4H_B$ )



# Plume Basics - Gas Thrust Region

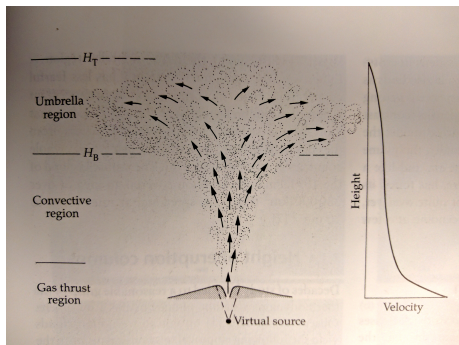
- magma vesiculates and fragments, accelerates as low viscosity gas+particle mix
- ejection velocity mostly governed by proportion of volatiles
- plinian eruptions of silicic magmas (4-5 wt% volatiles):  
>500 m/s (supersonic)



*Francis & Oppenheimer, 2004*

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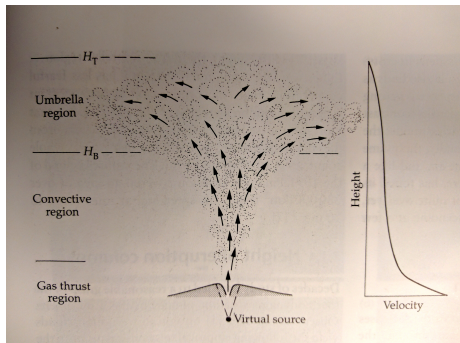
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- ejection velocity mostly governed by proportion of volatiles
- plinian eruptions of silicic magmas (4-5 wt% volatiles):  
>500 m/s (supersonic)
- denser than atmosphere but propelled by momentum
- starts to entrain ambient air:  
**gas thrust/jet phase**
- air heated by pyroclasts, reducing bulk density of plume
- up to a few km high



*Francis & Oppenheimer, 2004*

# Plume Basics - Convective Region

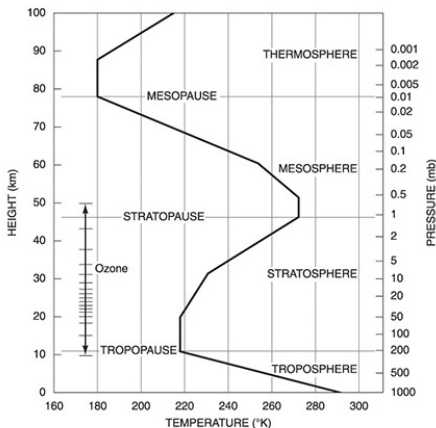
- hot gas is buoyant in atmosphere
- **convective region** owes buoyancy to heating of entrained air
- may lift plume several 10s of kilometers high



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# Plume Basics - Convective Region

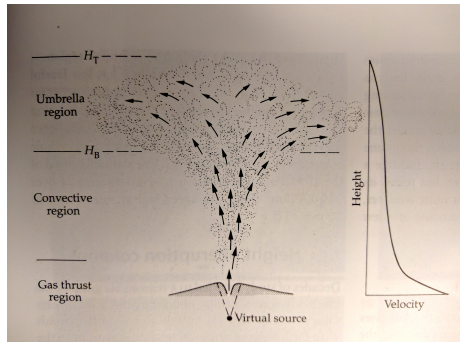
- hot gas is buoyant in atmosphere
- **convective region** owes buoyancy to heating of entrained air
- may lift plume several 10s of kilometers high
- can cross tropopause and puncture stratosphere
- ascent rates of 10s-100s m/s
- humid troposphere increases this:
  - condensation of water vapor yields latent heat
  - freezing of liquid water: more latent heat
- plume widens in cross section, density contrast lessens with altitude



Francis & Oppenheimer, 2004

# Plume Basics - Umbrella Region

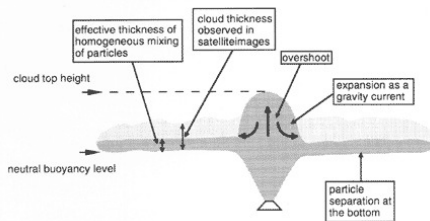
- plume will rise until reaching  $H_B$ : plume density == atmosphere density
- plume retains momentum, can rise a few kms more to  $H_T$



*Francis & Oppenheimer, 2004; Takehiro Koyaguchi;*

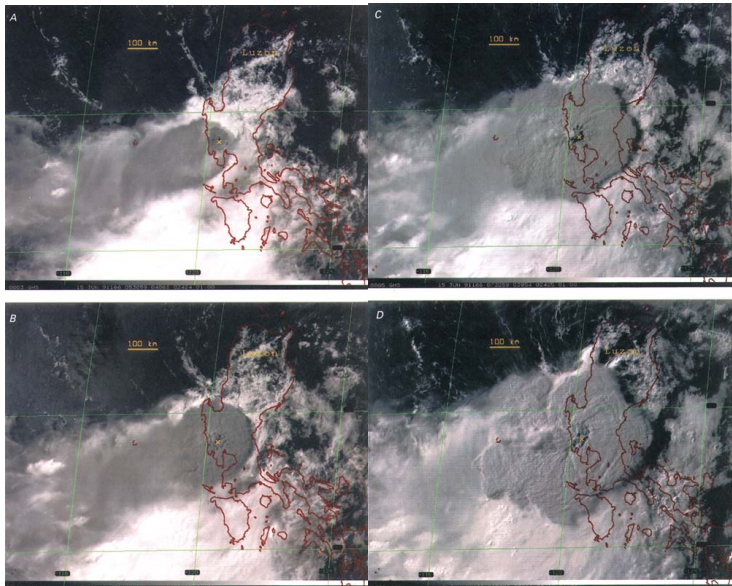
# Plume Basics - Umbrella Region

- plume will rise until reaching  $H_B$ : plume density == atmosphere density
- plume retains momentum, can rise a few kms more to  $H_T$
- horizontal spreading sets in, even upwind(!), gravity current driven
- most of tephra fallout from umbrella's base
- Pinatubo 1991 umbrella cloud covered area of  $230000 \text{ km}^2$



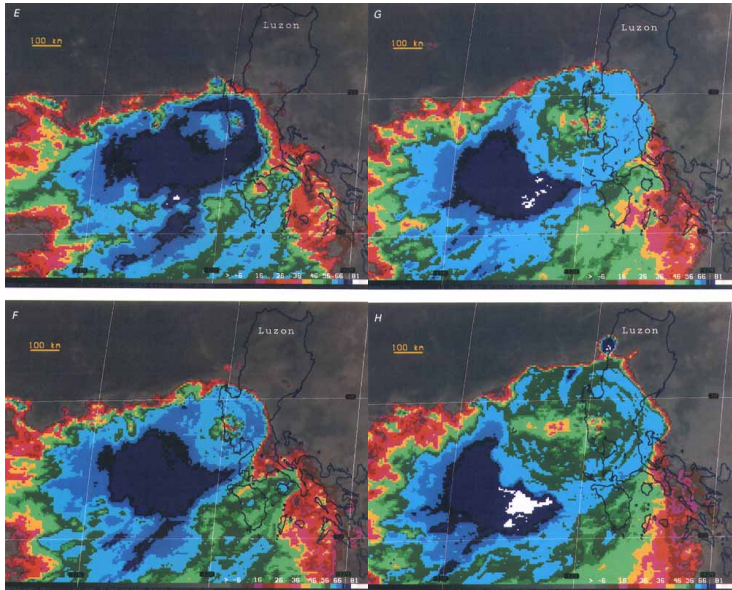
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# Plume Basics - Pinatubo



visible wavelengths from 13:40-16:40 local time *Self et al., Fire & Mud*

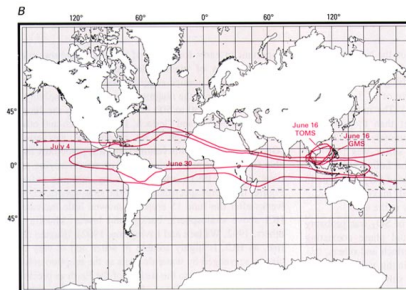
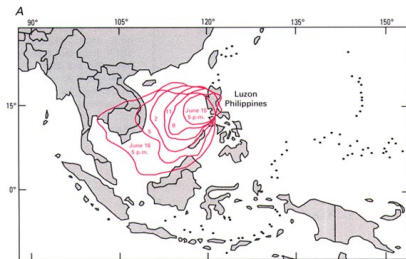
# Plume Basics - Pinatubo



Thermal-infrared wavelengths from 13:40-16:40 local time *Self et al., Fire & Mud*



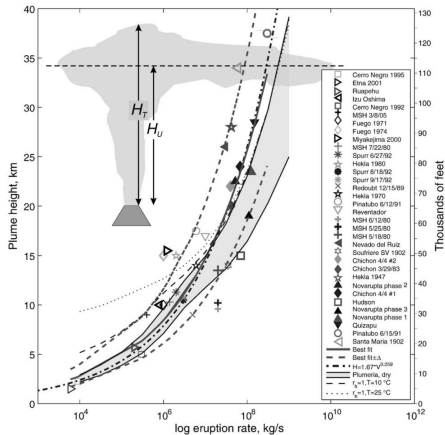
# Plume Basics - Pinatubo



Self et al., *Fire & Mud*

# Plume Basics - Plume Heights

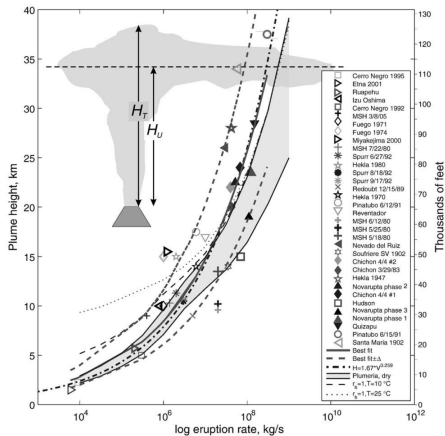
- relationship between height of a plume ( $H_T$  in km) and **dense rock equivalent** magma discharge rate ( $Q$  in  $m^3/s$ )



Mastin et al., 2009

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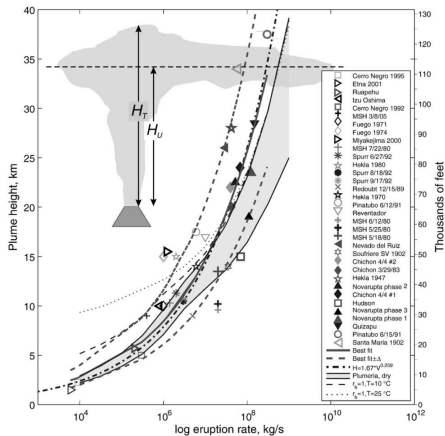
- relationship between height of a plume ( $H_T$  in km) and **dense rock equivalent** magma discharge rate ( $Q$  in  $m^3/s$ )
- $H_T = 2.00Q^{0.241}$  (Mastin et al, 2009) (converted mass eruption rate kg/s to discharge rate  $m^3/s$ )



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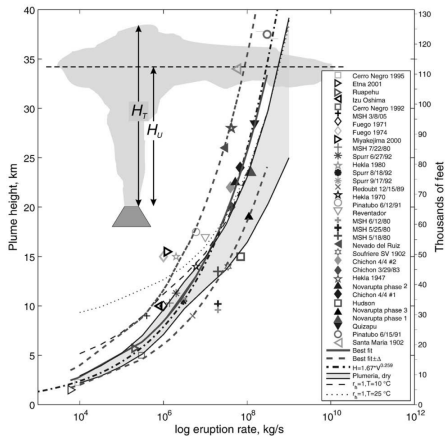
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- i.e. 16-fold increase in eruption rate required to double plume height
- tracking  $Q$  through eruption and integrating gives total erupted mass



Mastin et al., 2009

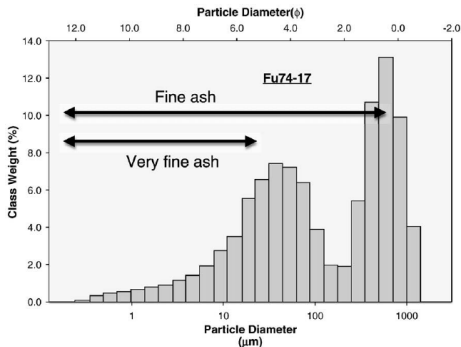
# Pyroclastic Fall Deposits

- deposits used to assign magnitude and intensity to prehistoric eruptions
- **deposit magnitude:** thickness or mass as function of distance from vent
- generally exponential thinning of tephra deposits from vent
- use to estimate total volume
- distal segment most difficult to assess (wide distribution, poorly preserved)



# Pyroclastic Fall Deposits

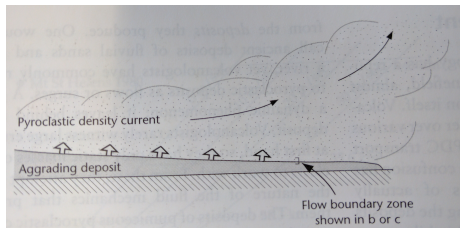
- **eruption intensity**: infer from relationships between grain size, grain density, column height, depositional characteristics
- large uncertainties introduced during field sampling (point samples)
- total grain size distribution (TGSD) is more important (input for ash dispersion models)
- TGSD difficult to measure:
  - deposits are widespread
  - grain size varies up section
  - distal deposits again poorly preserved
- critical data for source plume models



Fuego Volcano, *Rose & Durant, 2009*

# Pyroclastic Density Currents

- hot gravity driven currents
- travel at high velocities
- can form by lava dome or (sub- / ultra-) plinian column collapse
- high velocities, high temperatures make direct measurements impossible

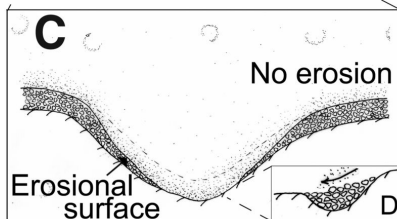
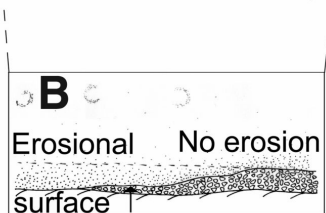
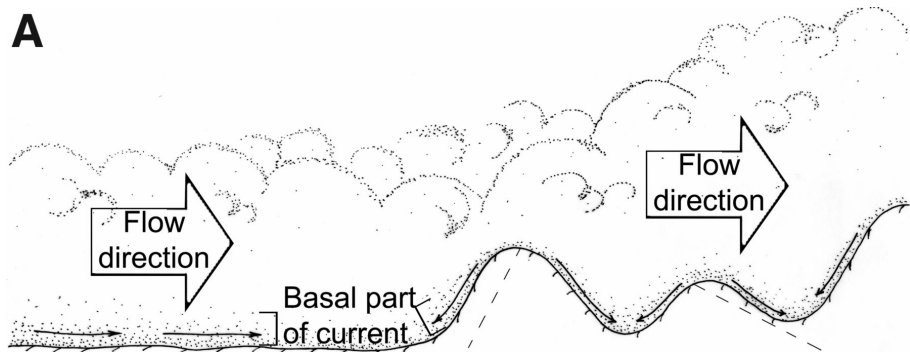


*Francis & Oppenheimer, 2004*



# Pyroclastic Density Currents

**A**



<http://geology.gsapubs.org/content/40/11/1035/F3.expansion.html>

<https://www.youtube.com/watch?v=Cvjwt9nnwXY>

# Volcanic Explosivity Index

- Devised by Chris Newhall & Stephen Self in 1982
- “Richter scale for eruptions” (note, we don’t use Richter scale anymore)
- based on volume of tephra, plume height
- often criticized / used for things it wasn’t made to do
- recently revised by Houghton and others (2013, *Geology*) to account for small explosive events

# Volcanic Explosivity Index

VEI	0	1	2	3	4	5	6	7	8
General Description	Non-Explosive	Small	Moderate	Moderate-Large	Large	Very Large			
Volume of Tephra (m <sup>3</sup> )		1x10 <sup>4</sup>	1x10 <sup>6</sup>	1x10 <sup>7</sup>	1x10 <sup>8</sup>	1x10 <sup>9</sup>	1x10 <sup>10</sup>	1x10 <sup>11</sup>	1x10 <sup>12</sup>
Cloud Column Height (km) Above crater Above sea level	<0.1	0.1-1	1-5	3-15	10-25			>25	
Qualitative Description	"Gentle,"	"Effusive"	← "Explosive" →		← "Cataclysmic," "paroxysmal," "colossal" →		← "Severe," "violent," "terrific" →		
Eruption Type (see fig. 7)	← Hawaiian →		← Strombolian →		← Vulcanian →		← Plinian →		← Ultra-Plinian →
Duration (continuous blast)	← <1 hr →		← 1-6 hrs →		← 6-12 hrs →		← >12 hrs →		
Maximum explosivity	Lava flow	← Phreatic →		← Explosion or Nuée ardente →		← Dome or mudflow →			
Tropospheric Injection	Negligible	Minor	Moderate	Substantial →					
Stratospheric Injection	None	None	None	Possible	Definite	Significant →			
Eruptions	976	1239	3808	1083	412	168	50	6	0

*Newhall & Self, 1982; Siebert et al., 2010*

# Volcanic Explosivity Index+

